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Swiss Confederation

Federal Department of Home Affairs FDHA Federal Office of Meteorology and Climatology MeteoSwiss



How to retrieve surface radiation and surface albedo from satellites?

Rebekka Posselt, Aku Riihelä With support from: Richard Müller, Jörg Trentmann





PART I

Solar radiation (SIS, SID)

- MagicSol Retrieval for historical radiation datasets
- LookUpTable radiation retrieval

Longwave radiation (SDL)

AVHRR-CLARA (GAC) algorithm (very short)

PART II

Surface albedo (SAL)

Retrieval Overview







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Part I Surface radiation retrievals

Rebekka Posselt (MeteoSwiss)

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PART II Surface albedo (SAL)



- Historical = Meteosat 2 7 (1983 2005)
 - = MVIRI instrument
- Benefits:

U

- 23 years of high resolution data \rightarrow climatology
- Challenges:
 - Only three available channels (VIS, IR, WV)
 - Satellite operations not designed for climate studies
 - Different satellites (inhomogeneities)
 - Poorly documented (M2-4 \rightarrow missing calibration)



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 - = MVIRI instrument
- Benefits: •

J

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 - Only three available channels (VIS, IR, WV)
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 - Different sa Self-calibration g calibration)



• Retrieval scheme

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- 1. Get cloud information
 - "Effective Cloud Albedo" (*CAL*)
 - From satellite
- 2. Get clear sky information
 - "Clear Sky Radiation" (*Rad_{cs}*)
 - From LookUpTables
- 3. Combine 1. & 2.

$$Rad = f(CAL) \cdot Rad_{cs} \approx (1 - CAL) \cdot Rad_{cs}$$

$$\rightarrow$$
 Rad = SIS or SID





- 1. Get cloud information
 - Get clear sky = cloud free = surface only image from a series of original images (usually the darkest pixel in the series)







1. Get cloud information

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Get Selfcalibration via p_{max} = 95% percentile of all counts in target region (South Atlantic)





1. Get cloud information

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 All information together give the "effective cloud albedo" (CAL, a.k.a. cloud index) → Heliosat method





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 All information together give the "effective cloud albedo" (CAL, a.k.a. cloud index) → Heliosat method

$$CAL = \frac{\rho - \rho_{min}}{\rho_{max} - \rho_{min}} \rightarrow Cloud \text{ image} \\ \sim Maximum range of pixel brightnesses}$$

- Overcast \rightarrow *CAL=*?
- Clear-sky \rightarrow CAL = ?
- Fresh snow \rightarrow CAL < 0, 0<CAL<1, CAL>1



LookUpTables •

Surface Albedo

Climatology

 \rightarrow Fast

Ozone

MagicSol

 \rightarrow Obtained from a Radiative Transfer Model

Müller et al. (2009)

http://sourceforge.net/projects/gnu-magic

Rad_{cs}





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PART II Surface albedo (SAL)



• Used for

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- Operational (= regularly updated) radiation products
- AVHRR-CLARA (GAC) radiation dataset
- Benefits:
 - Physical approach
 - Applicable to geostationary <u>and</u> polar orbiting satellites
- Challenges:
 - Multispectral information required for cloud detection
 - Auxiliary data required (e.g., surface albedo)



- Retrieval scheme
 - 1. Cloud detection
 - 2. <u>Cloud free:</u> use clear-sky gnu-magic (see MAGICSOL)



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- Retrieval scheme
 - 1. Cloud detection
 - 2. <u>Cloudy:</u> Get atmospheric trans-missivity τ from LUT
 - Use satellite and model data as input
 - 3. Calculate Rad

$$Rad = E_0 \cdot cos\Theta_z \cdot \tau$$

- E_0 = solar constant = 1362 Wm⁻²
- Θ_z = sun-zenith angle
- τ = atmospheric transmissivity







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$$Rad = E_0 \cdot cos\Theta_z \cdot \tau$$



- $E_0 = \text{solar constant} = 1362 \text{ Wm}^{-2} \leftarrow \text{k}$
- Θ_z = sun-zenith angle
- τ = atmospheric transmissivity
- ← known
 - ← known
 - ← use cloud-magic



0

1. Get atmospheric transmissivity τ (cloud-magic)



Müller et al. (2009)

http://sourceforge.net/projects/gnu-magic



- E_0 = solar constant = 1362 Wm⁻²
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LookUpTable radiation retrieval

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Arctic-SAL



AVHRR-CLARA (GAC) SDL retrieval - very short

- GAC = "global area coverage"
 - = AVHRR instrument (1982-present), polar orbiting
- SDL mainly determined by Temperature and humidity close to the earth's surface
 - Cannot be observed by satellites
 - → all SDL products from satellites need additional data (e.g., reanalysis, NWP)
- CM SAF GAC SDL uses ERA Interim SDL as basis
- Cloud information of GAC are used to refine ERA SDL





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PART II SAL retrieval algorithm

Aku Riihelä FMI





Surface albedo

Radiation budget at surface: $E_{net} = SW \downarrow - \alpha * SW \downarrow + LW \downarrow - LW \uparrow$

The resulting net energy is available for surface heating, snow melt, heat fluxes etc.



Mean surface albedo of the Arctic, June 2009



Mean surface albedo of the Arctic, August 2009



100 90 80



The SAL algorithm

- A shortwave black-sky surface albedo product
 - Black-sky = direct solar flux only, all atmospheric effects removed
- A radiometric and geolocation correction for topography effects on AVHRR images
- Dedicated algorithms for vegetated surfaces, snow/ice, and water
- Atmospheric correction with SMAC
- BRDF correction over vegetated surfaces





0. Topography correction



In the first part, we correct the geolocation of the AVHRR pixels for the true terrain height effects using a global DEM



In the second part, we correct the observed reflectances for effects caused by the various slopes and shadowed areas in an AVHRR pixel



1. Atmospheric correction

- Atmospheric effects need to be removed from the observed TOA reflectances
- We use the Simplified Model for Atmospheric Correction (SMAC) [Rahman & Dedieu, 1994].
- Required Inputs
 - Visible + near IR TOA reflectances
 - Aerosol Optical Depth (AOD) content of the atmosphere (set <u>constant to</u> 0.1)
 - Total ozone column (O3) (constant at 0.35 (atm cm))
 - Total column water vapour and surface pressure (taken from atmospheric model, ECMWF / DWD (g/cm^2))



2. BRDF correction

- Applied the model of Roujean (1992) with an update by Wu et al. (1995).
- The model considers the bidirectional reflectance of a surface to consist of three "kernels":





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k terms describe the reflectance contributions from:

- nadir-viewing & overhead Sun situation (k0),
- geometric and volume scattering terms k1 and k2.

The *f* terms describe the dependency of the model from the viewing/illumination geometry of the scene.



2.5. Anisotropy sampling of snow

- The reflectance anisotropy properties of snow vary widely with snow type!
 - Very difficult to model universally without universal data on snow physical characteristics
- Our solution: sample the anisotropy directly and consider the mean of the samples to represent the albedo.
- The strategy works if we have enough samples of the BRDF...which fortunately is the case when using AVHRR in the high latitudes (where snow exists)!





Reflectance sampling distribution at Summit Camp, Greenland Ice Sheet, summer 2005



3. Narrow-to-broadband conversion

Satellite imagers cover only a part of the solar spectrum – algorithms needed to convert observed (spectral) albedo to full broadband albedo!



NTBC algorithms separated by instrument (SEVIRI / AVHRR) and land cover (vegetation, snow, water)

Vegetation-AVHRR: Liang (2000)

Vegetation-SEVIRI: Van Leeuwen & Roujean (2002)

Snow: Xiong et al. (2002)

Water (LUT-based): Jin et al. (2004)



When all is said and done...

- We have retrieved a broadband black-sky surface albedo for a satellite image*
- The instantaneous images are then projected into a common map grid and averaged over a pentad/week/month to create the product we distribute to You, the user.



* Multiple images required for a robust snow albedo retrieval



Limitations of the algorithm

- 1. Sun Zenith Angle of the scene has to be less than 70 degrees and the Viewing Zenith Angle (of the satellite) less than 60 degrees
 - Retrievals would be unreliable outside these bounds
- 2. Cloud masking errors do occur sporadically
 - Cloud reflectance propagates into an albedo "retrieval"
- 3. Aerosols and O3 concentrations currently constant in retrievals
 - Increased uncertainty over areas where AOD is high (see figure below)
 - O3 effect is much smaller than the aerosol effect
- 4. Coarse resolution (15 km², 0.25 degrees, 25 km²) may not allow for accurate small-scale studies

Problems using SAL? You can contact me at aku.riihela@fmi.fi



Thanks for tuning in!

