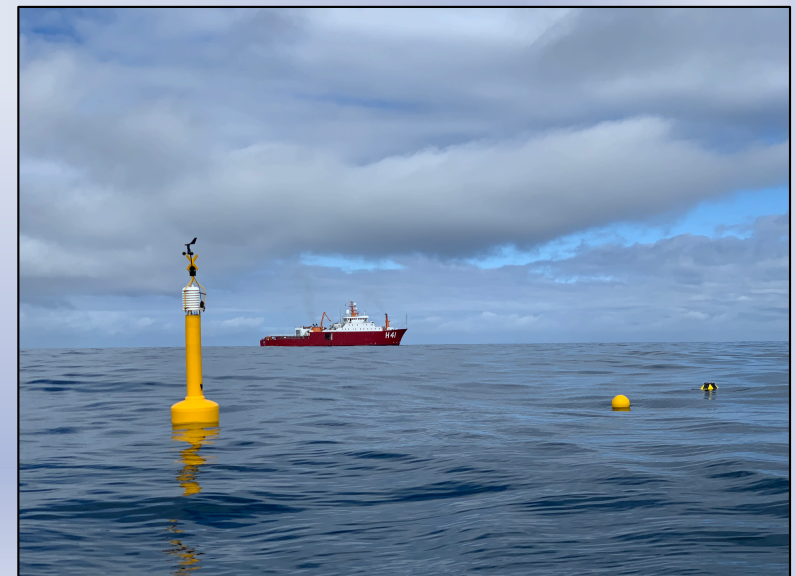
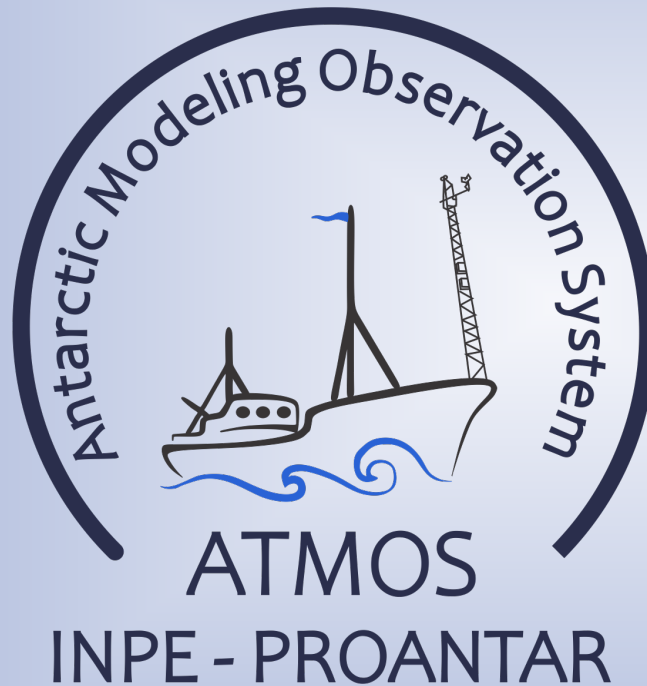
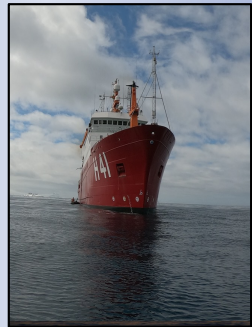


Mesoscale oceanic eddy-induced modifications to **air-sea heat** and CO₂ fluxes



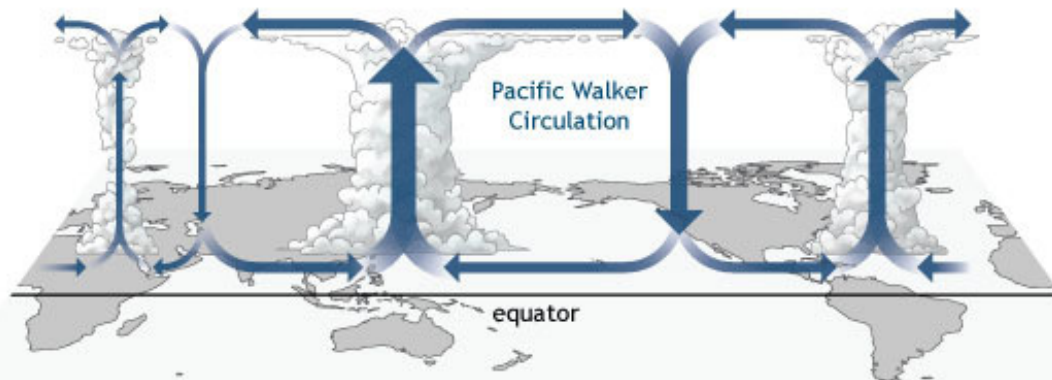
Dr. Luciano Ponzi Pezzi

Laboratory of Ocean and Atmosphere Studies (LOA)
Earth Observation and Geoinformatics Division (OBT)
National Institute for Space Research (INPE)



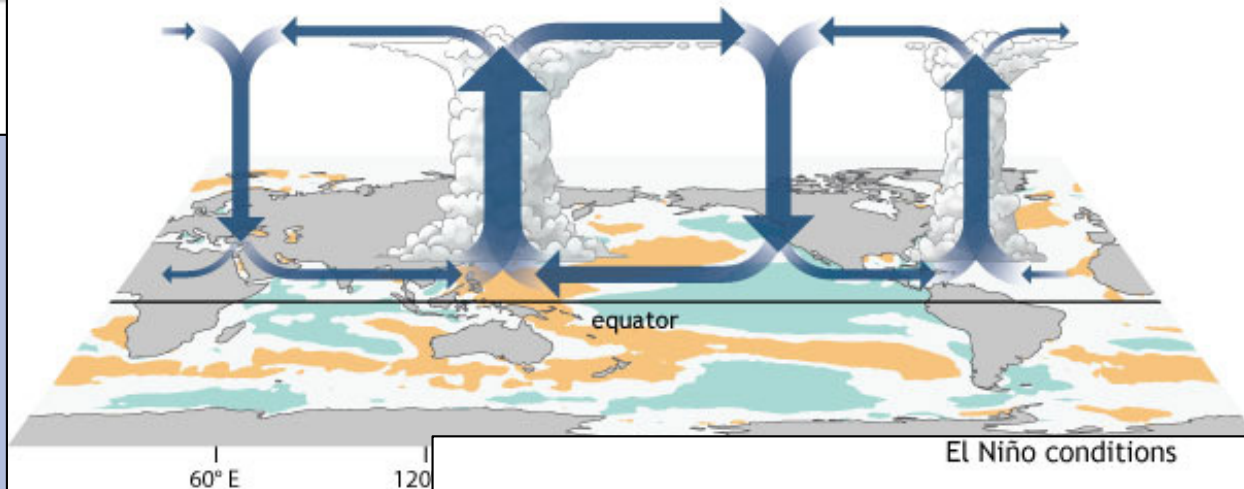
EUMETSAT-SOLAS
December - 2021

Neutral conditions

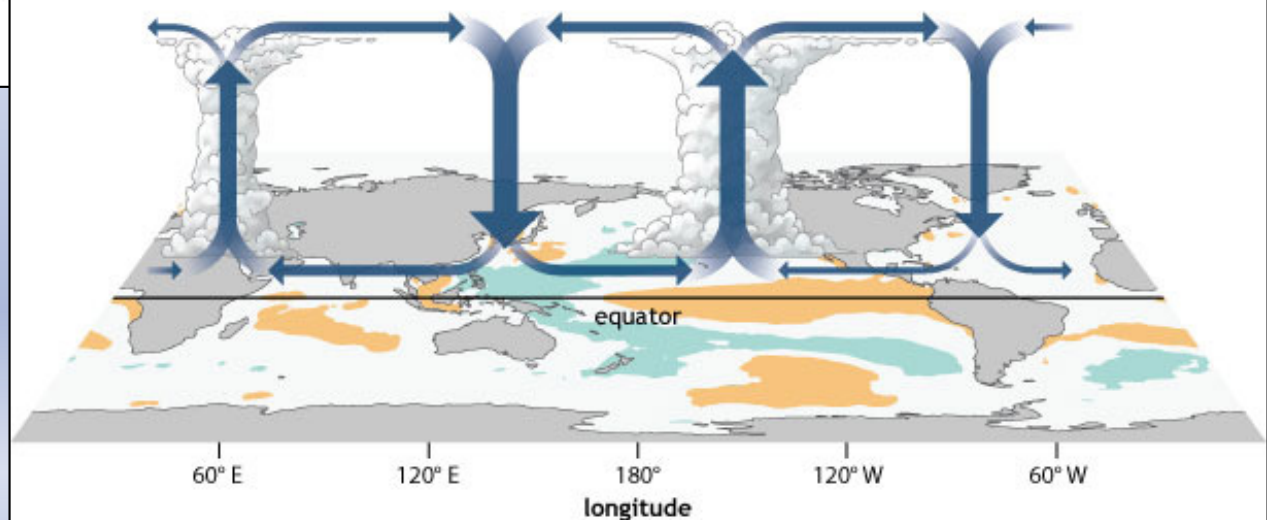


When it comes to O-A interaction
Then one thinks about basin scale
phenomena with global impacts,
teleconnections, monthly,
seasonal, decennial scales, etc..

La Niña conditions



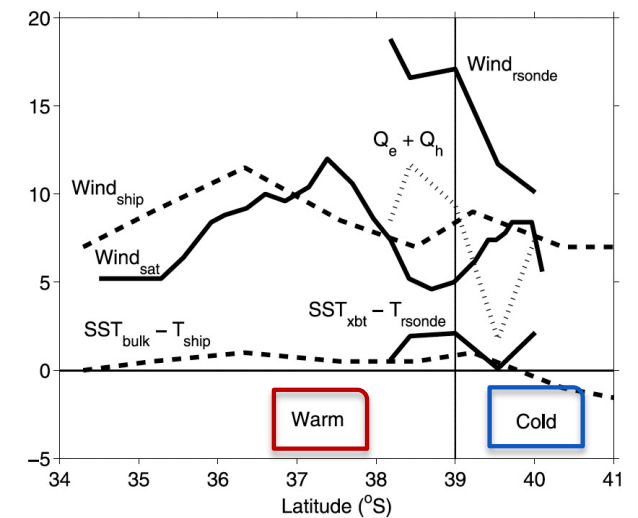
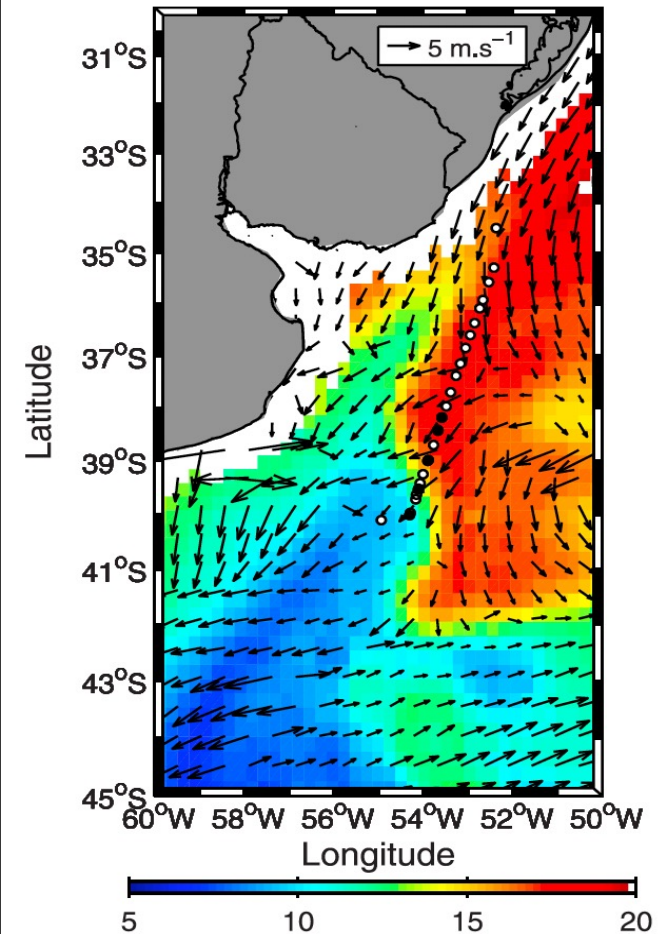
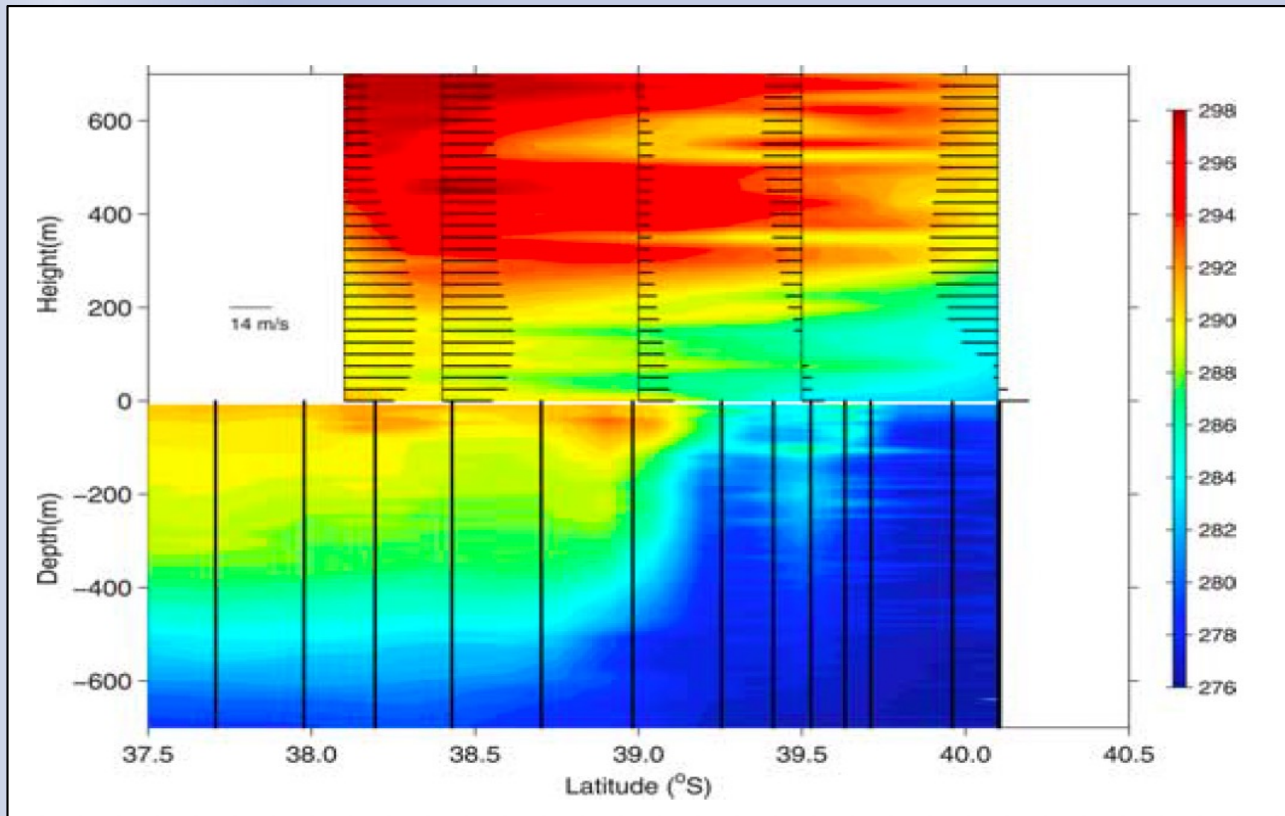
El Niño conditions



ENSO - (El Niño, La Niña)
Atlantic Dipole (or Variability) (TAV)
Pacific Decadal Oscillation (PDO)
and many
others
large scale variability phenomena

Ocean-atmosphere in situ observations at the Brazil-Malvinas Confluence region

L. P. Pezzi,¹ R. B. Souza,² M. S. Dourado,² C. A. E. Garcia,³ M. M. Mata,³
and M. A. F. Silva-Dias¹



This was our first O-A cruise and
we got a “textbook case”
In situ observations



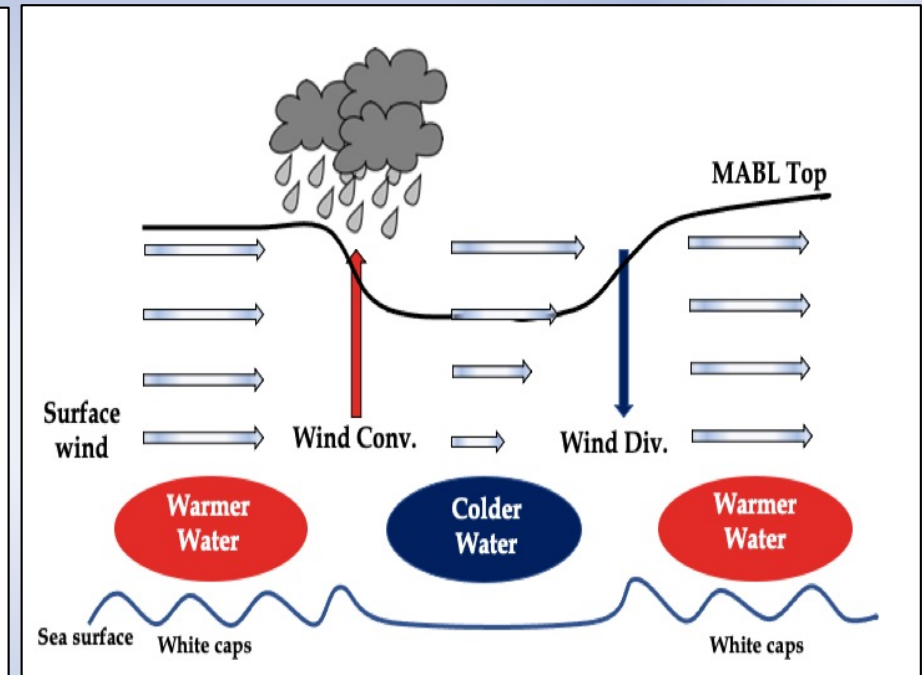
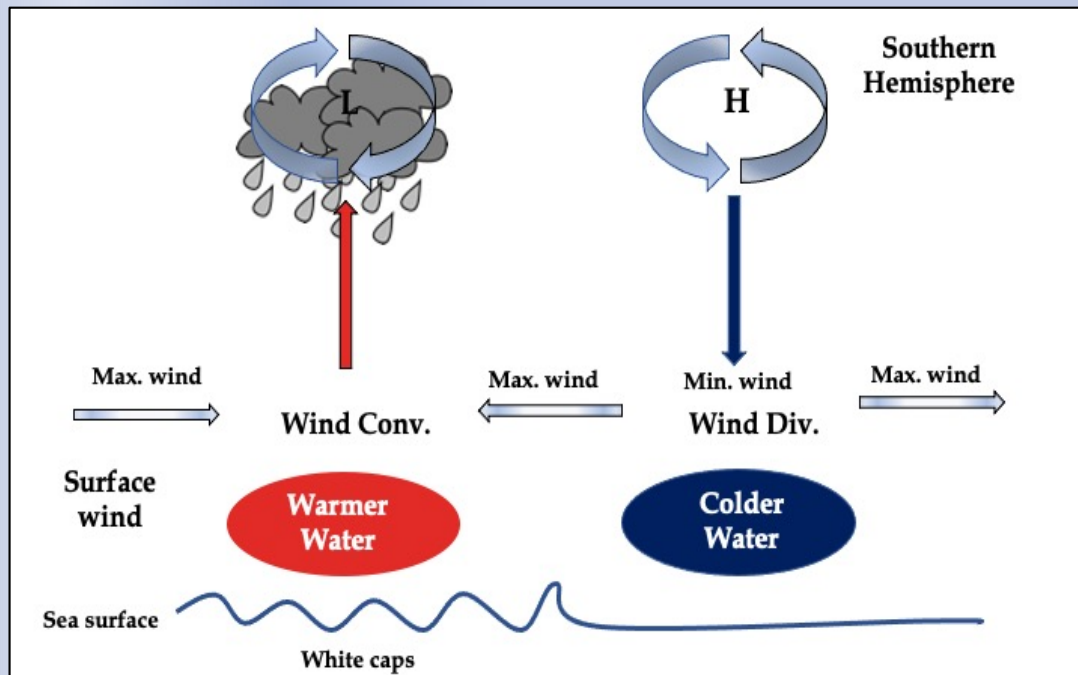
O-A interactions in strong SST gradient regions



TWO POSSIBLE PHYSICAL STABILITY MECHANISMS OF CLAM:

Hydrostatic Stability

Static Stability



Lindzen and Nigan (1987), Wallace *et al.* (1989) surface wind is affected by pressure gradient at sea level (SLP). *Hydrostatic stability*.

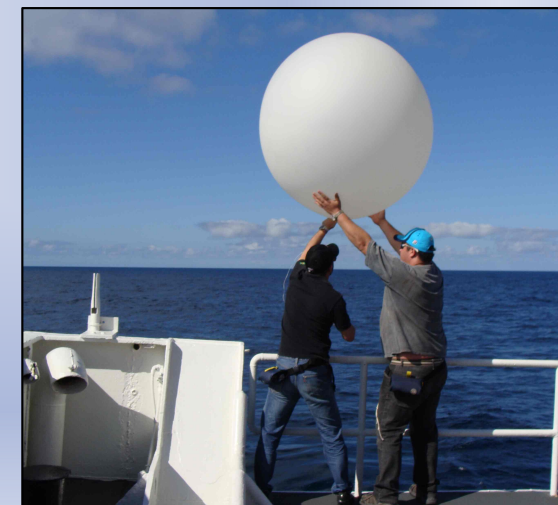
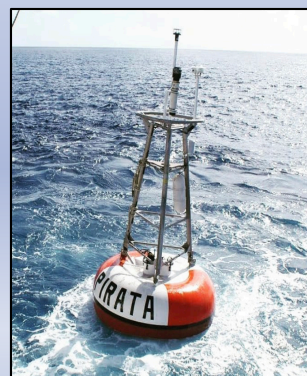
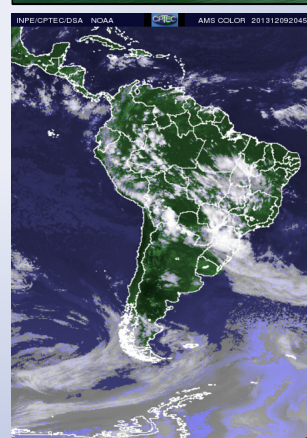
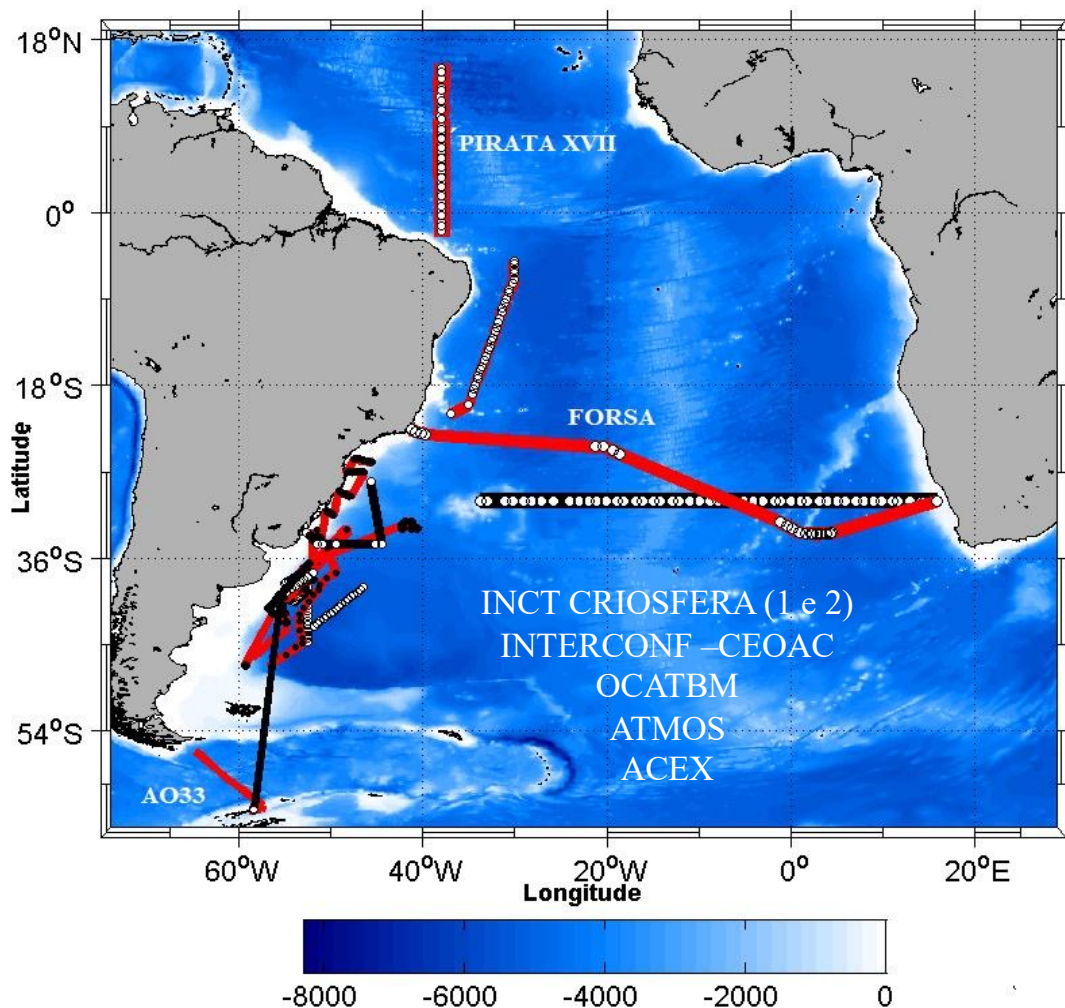
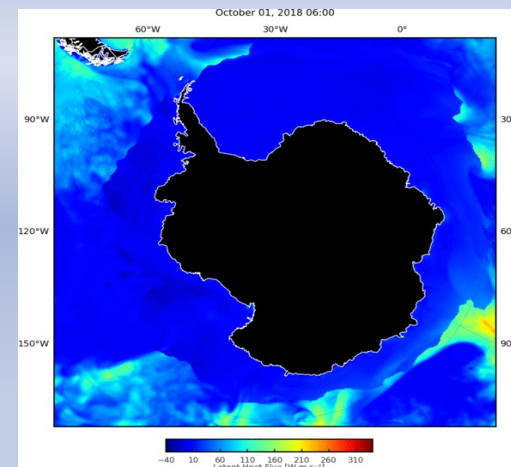
Hayes *et al* (1989) - surface wind is affected by the turbulence of the atmospheric boundary layer. *Static stability*





How do we study air-sea interaction processes?

- *In situ* observations (opportunity x permanent)
- Satellite data
- Numerical Modeling (global, regional, coupled)



EUMETSAT-SOLAS
December - 2021

INTERCONF
OPERANTAR 32
14 to 20 October 2013



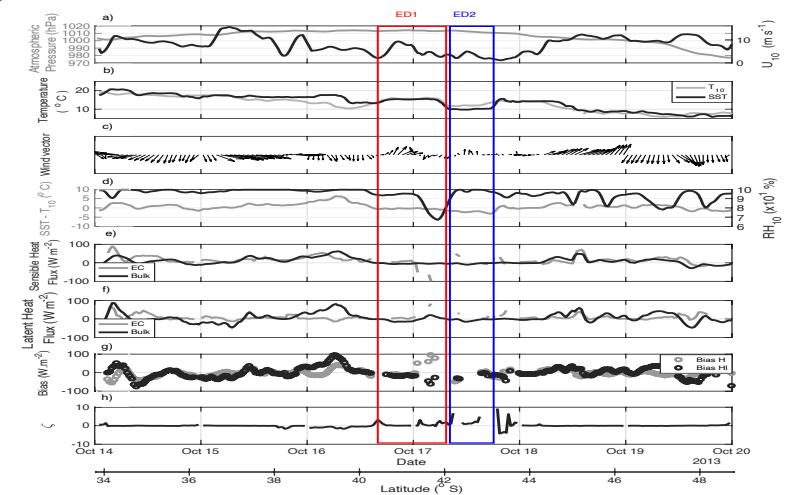
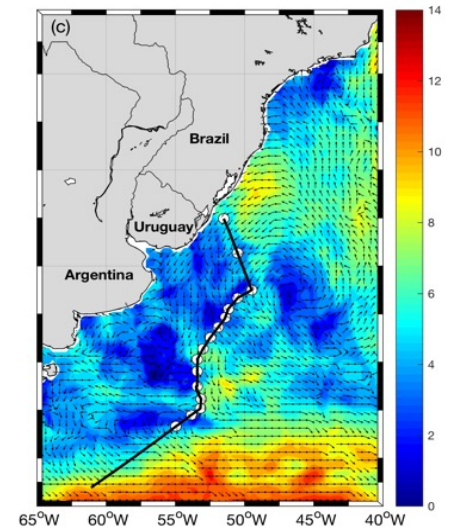
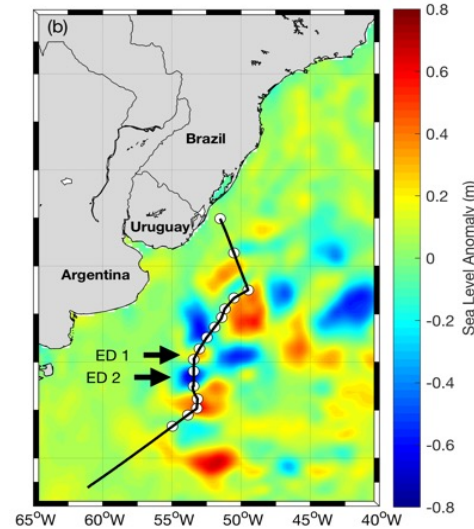
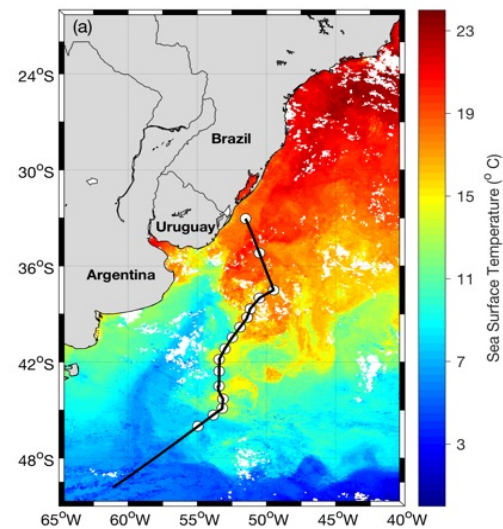
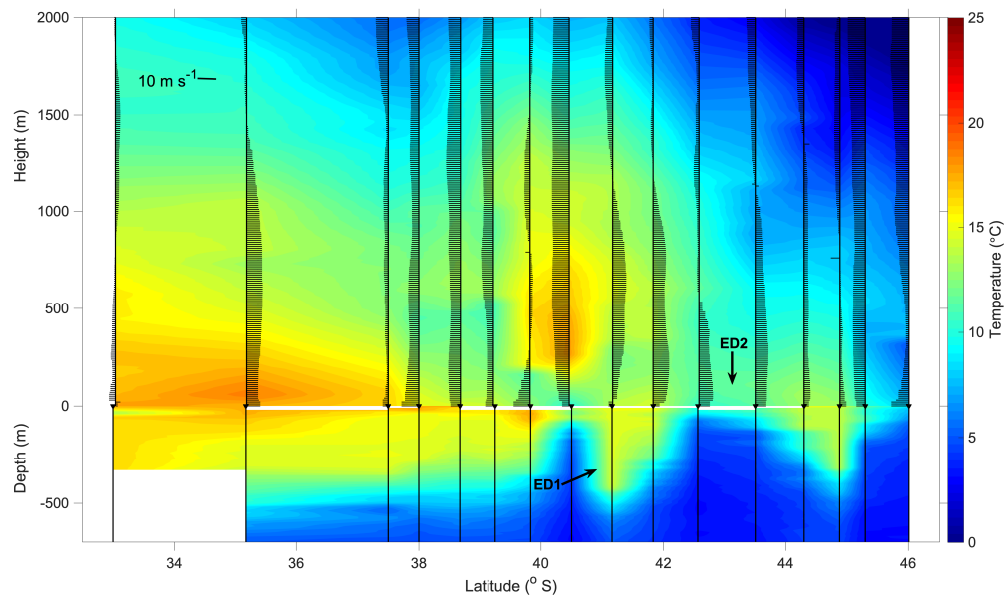
remote sensing

<https://www.mdpi.com/journal/remotesensing>

Article

Air-Sea Interactions over Eddies in the Brazil-Malvinas Confluence

Ronald Souza ^{1,*}, Luciano Pezzi ², Sebastiaan Swart ^{3,4} , Fabrício Oliveira ⁵ and Marcelo Santini ²



Increase – ED1:
78% sensible and 55% in latent heat fluxes

Decrease – ED2:
49% sensible and 25% in latent heat fluxes

EUMETSAT-SOLAS
December - 2021

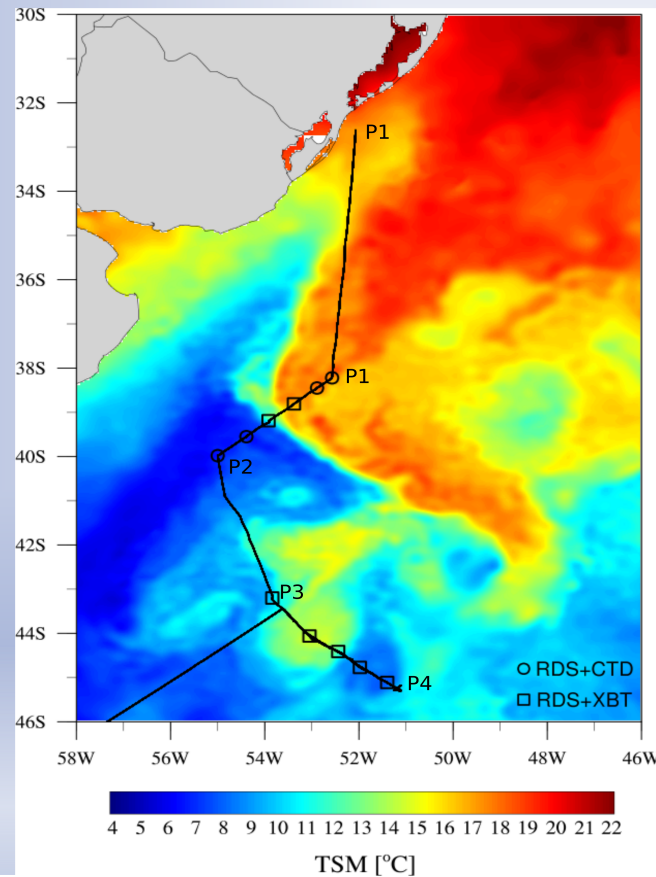
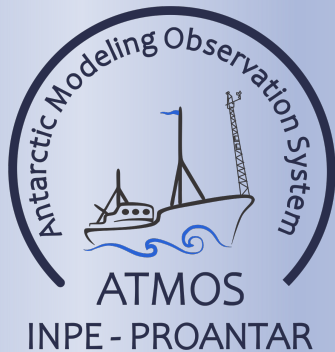


OPEN

Oceanic eddy-induced modifications to air-sea heat and CO₂ fluxes in the Brazil-Malvinas Confluence

Luciano P. Pezzi^{1✉}, Ronald B. de Souza², Marcelo F. Santini¹, Arthur J. Miller³, Jonas T. Carvalho¹, Claudia K. Parise⁴, Mario F. Quadro⁵, Eliana B. Rosa¹, Flavio Justino⁶, Ueslei A. Sutil¹, Mylene J. Cabrera¹, Alexander V. Babanin⁷, Joey Voermans⁷, Ernani L. Nascimento⁸, Rita C. M. Alves⁹, Gabriel B. Munchow⁹ & Joel Rubert¹⁰

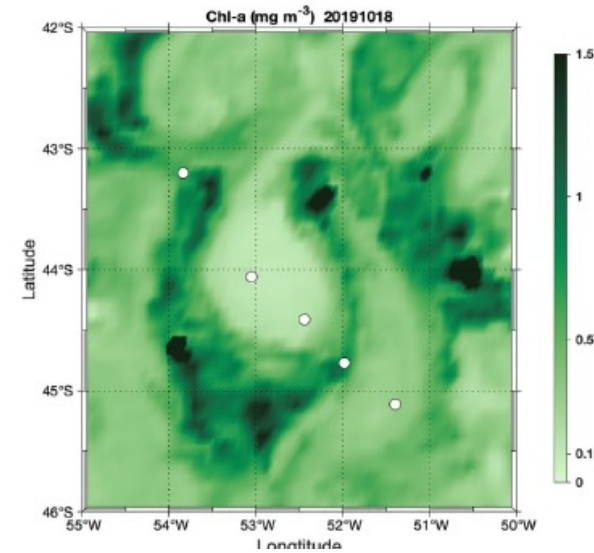
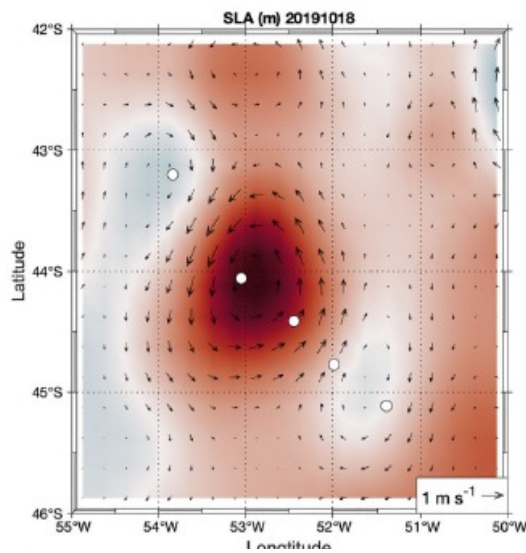
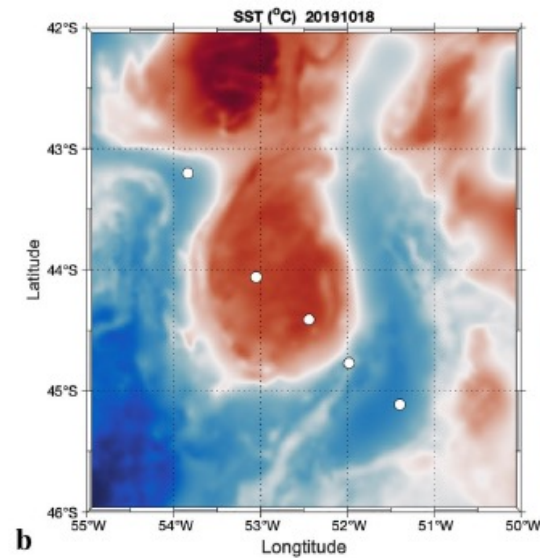
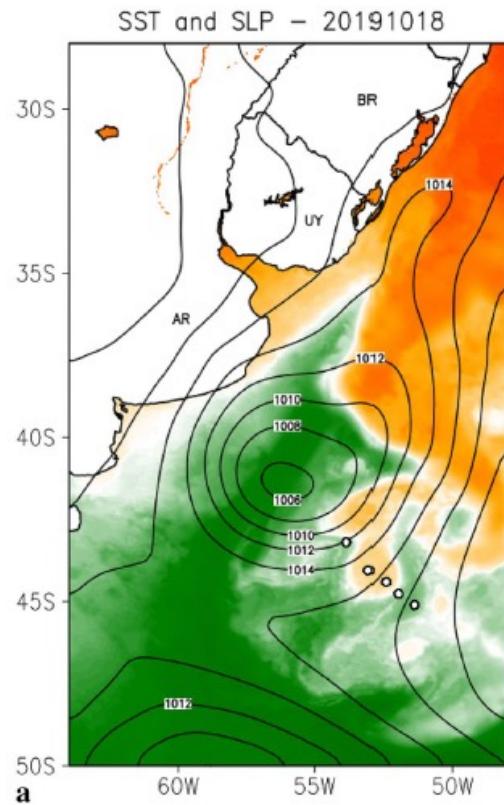
ATMOS Project
OPERANTAR 38
October/November 2019

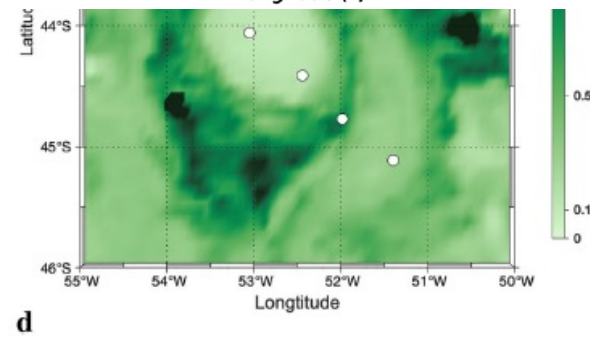
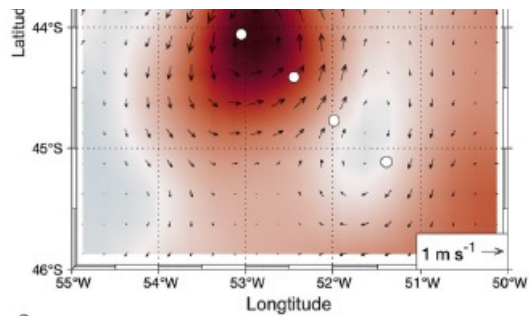
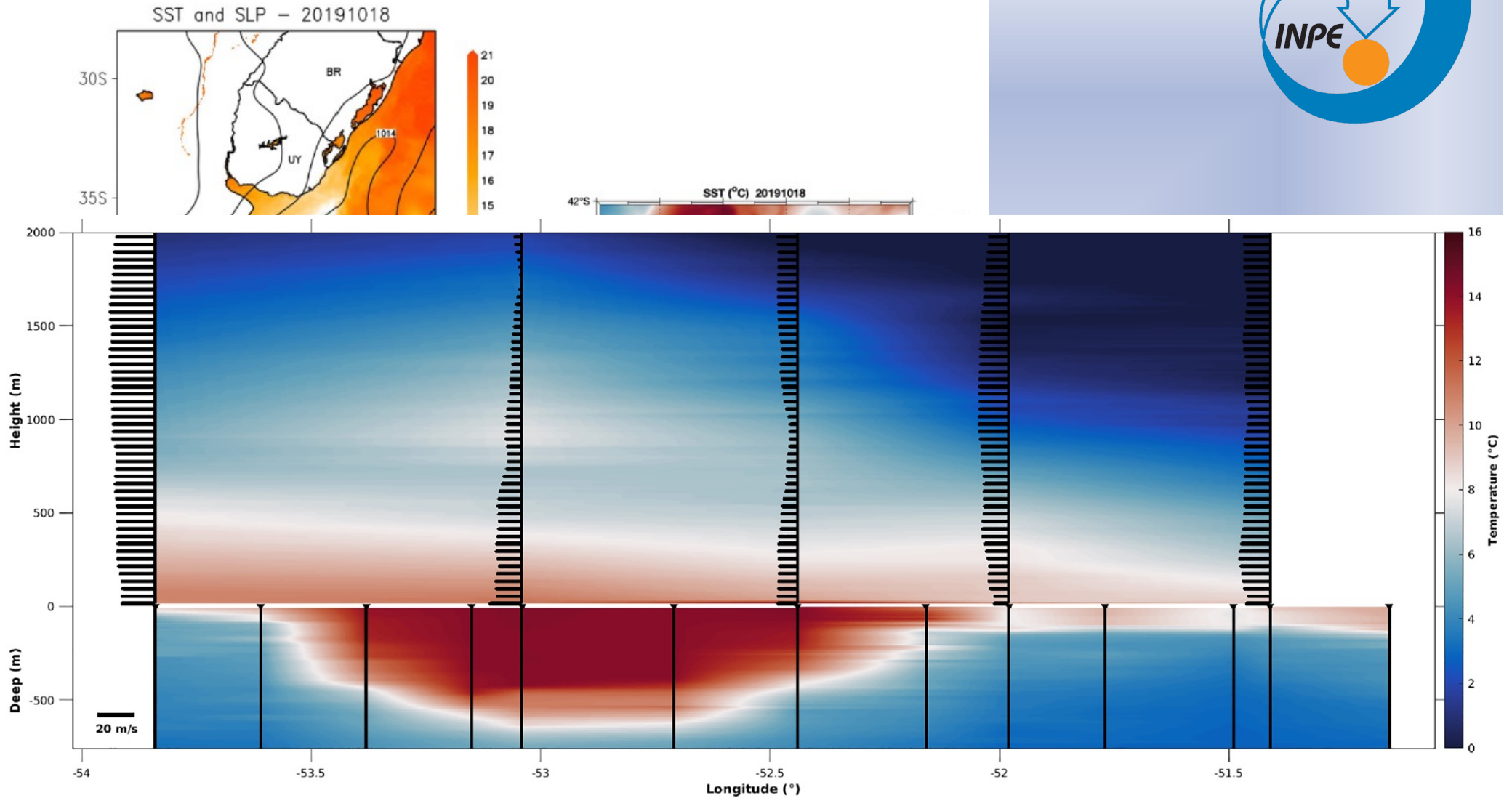


EUMETSAT-SOLAS
December - 2021



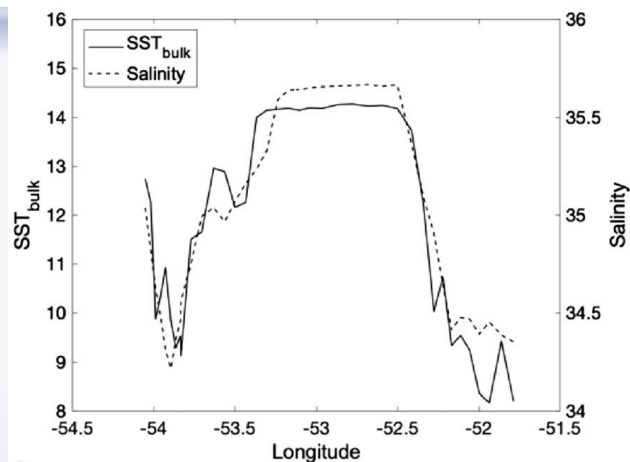
Three different satellite products



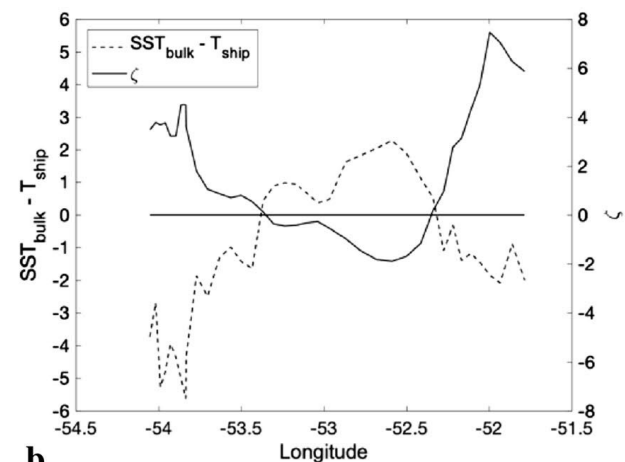


EUMETSAT-SOLAS
December - 2021

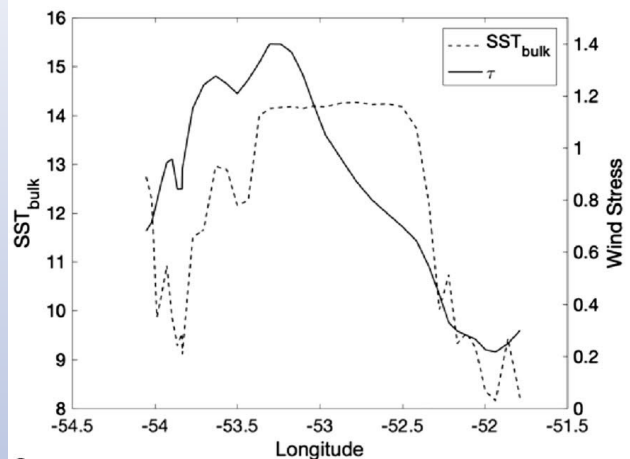
Look at the power of the warm core eddy in OP38



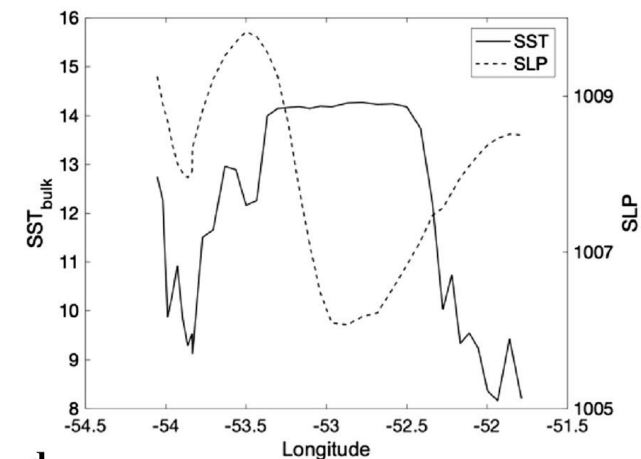
a



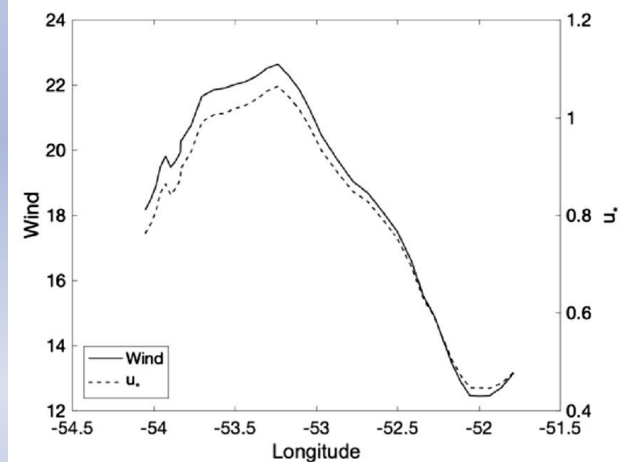
b



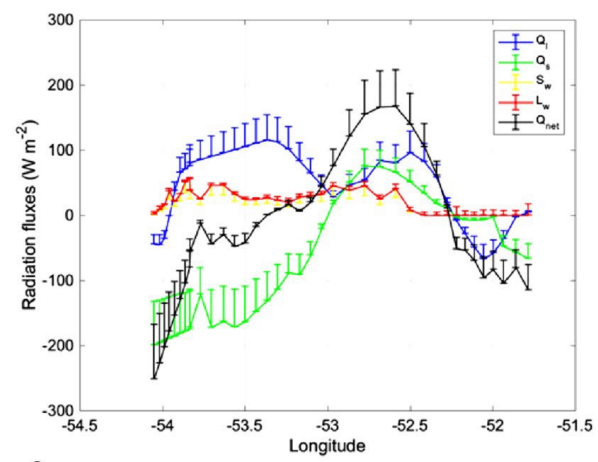
c



d



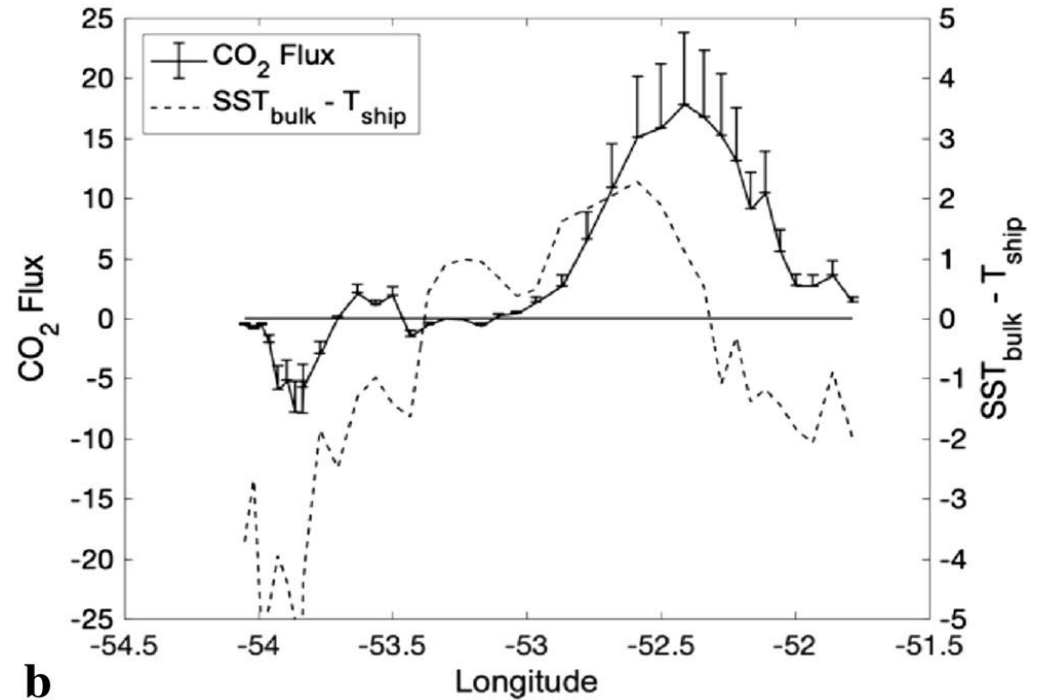
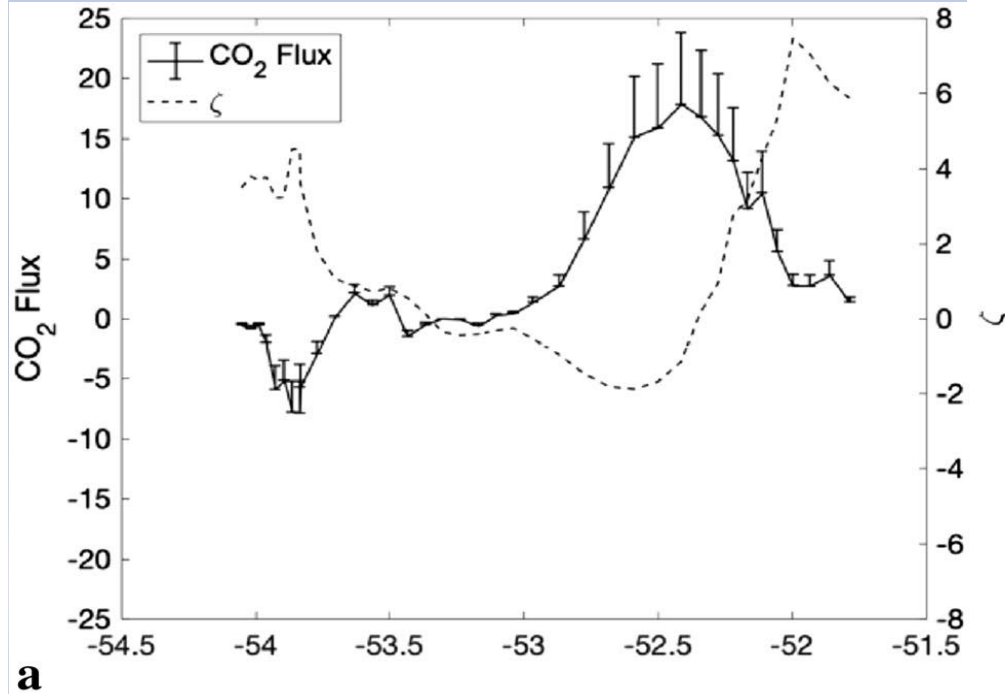
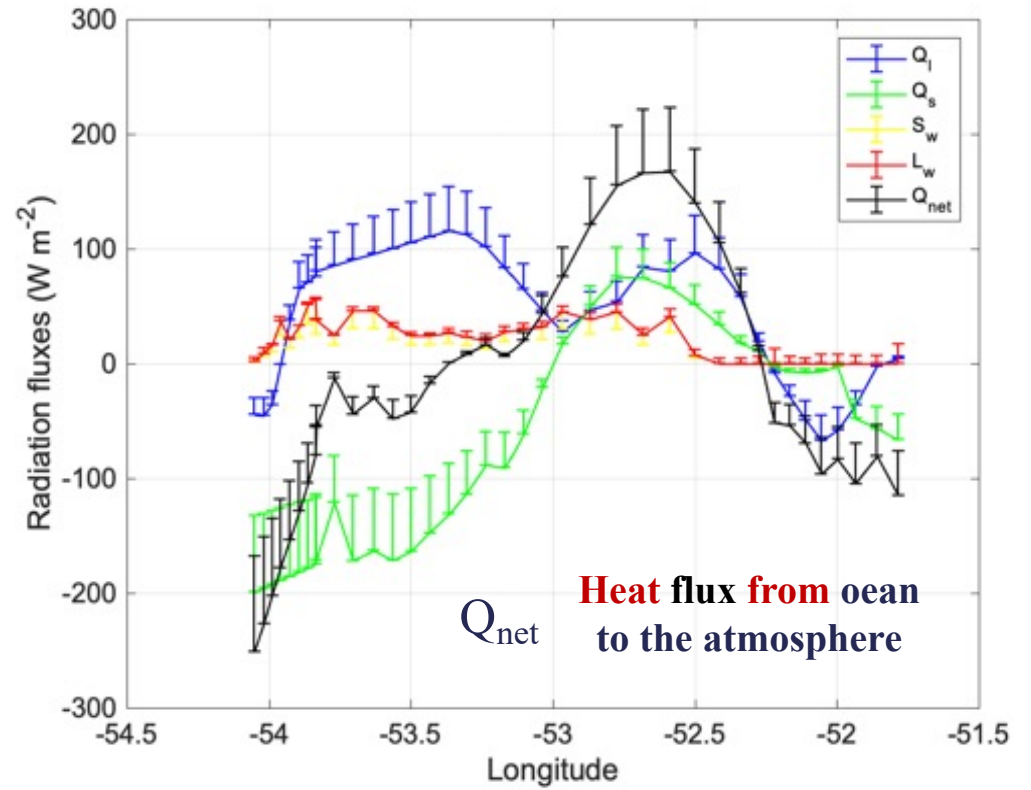
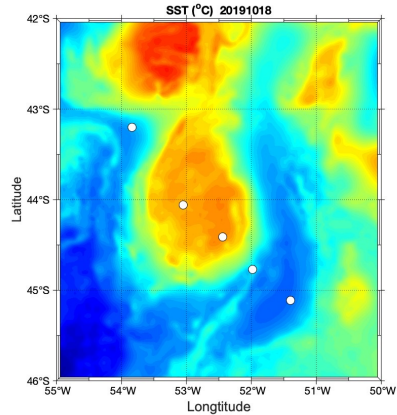
e



f



Look at the power of the warm core eddy in OP38



In addition to getting turbulent fluxes
observations...
we want to try our own parameterizations...

$$FCO_{2BK} = s.k.\Delta pCO_{2(mar-ar)}$$

CO₂ fluxes

$$\frac{\tau}{\rho} = \overline{u'w'} = u_*^2 = C_{D10}\overline{u_{10}}^2.$$

Momentum fluxes

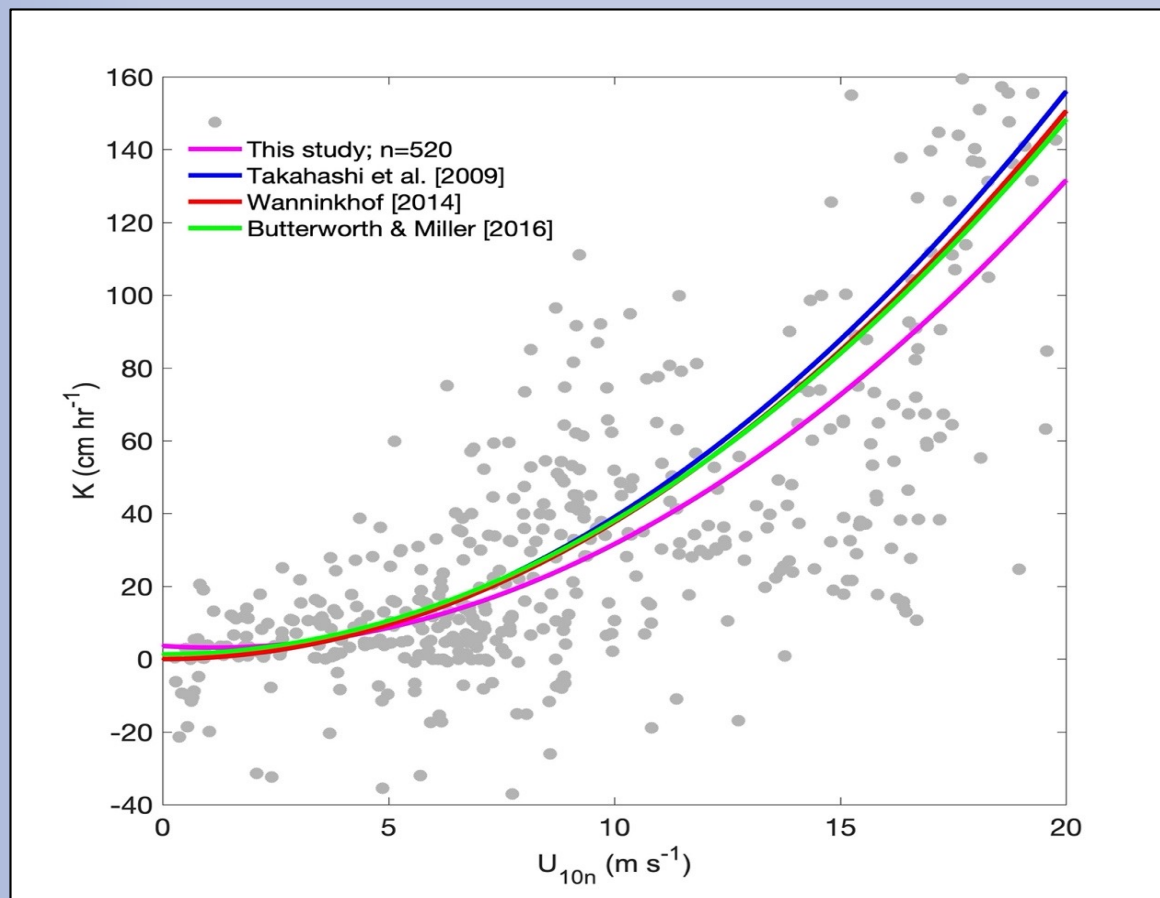
analogue

$$C_{DN10} = (0.75 + 0.067\overline{u_{N10}}) \times 10^{-3}.$$

Heat fluxes

both the **transfer velocity coefficient**
and **drag coefficient** ...
are bulk "**physical-statistical**" parameterizations

In addition to getting observational
turbulent fluxes ...
we want to try our own parameterizations...



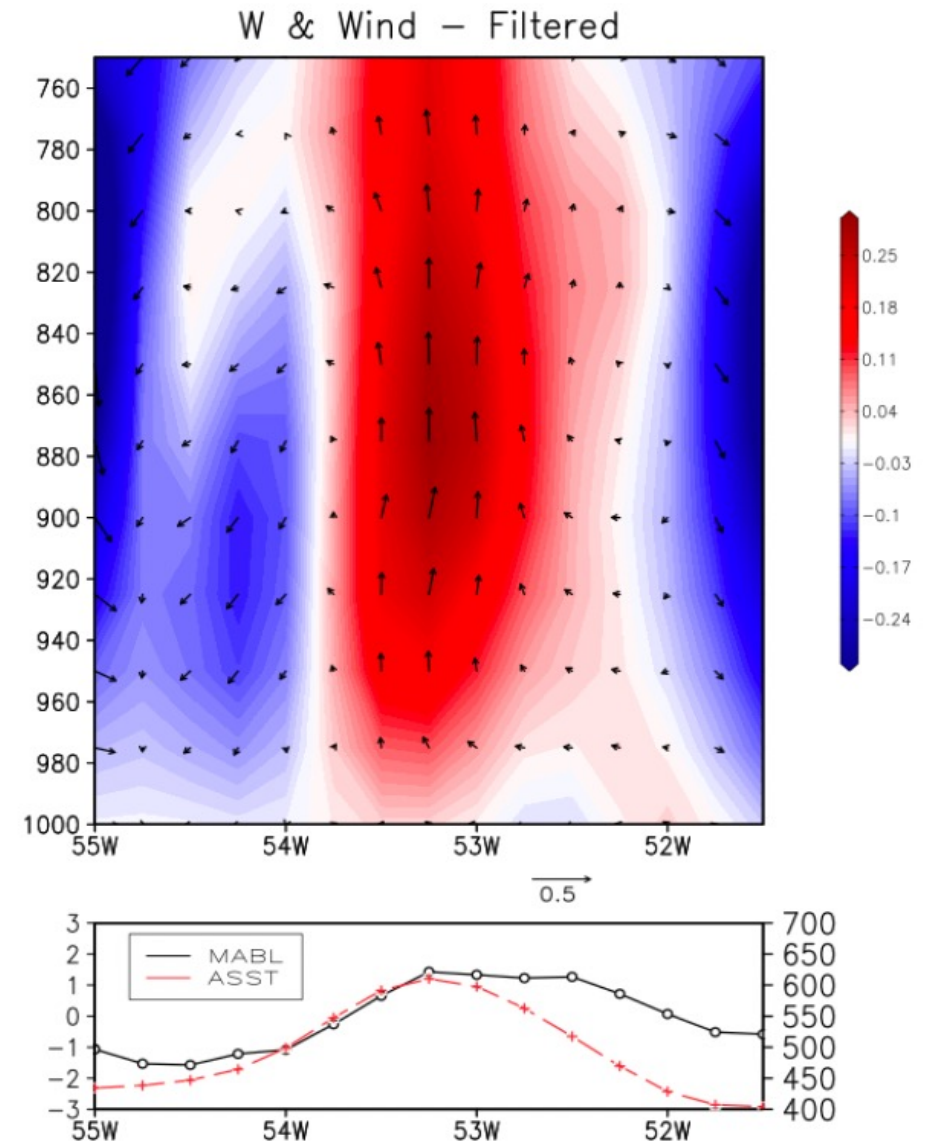
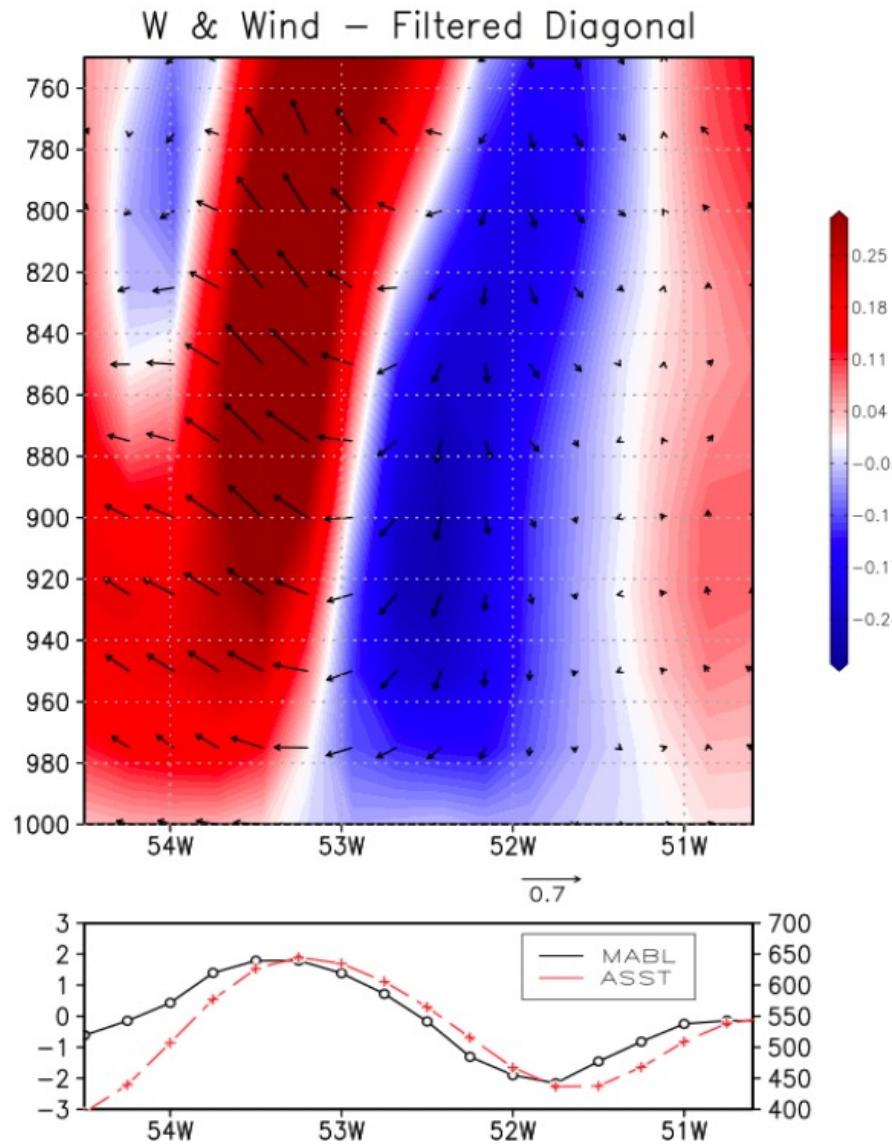
“transfer velocity
coefficient (K)” ...
is a bulk
“physical-statistical”
parameterization

$$FCO_{2BK} = s.k. \Delta p CO_{2(sea-air)}$$

This is our first try on
parameterization to SWA

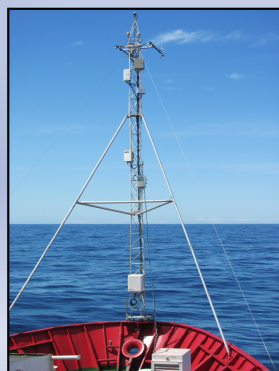
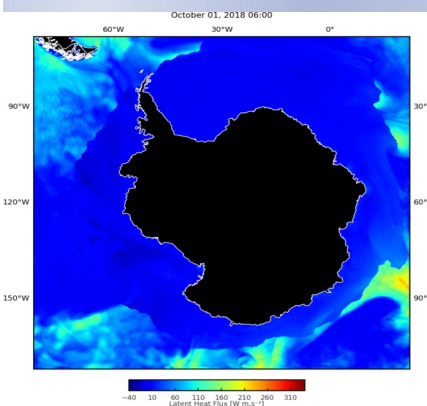
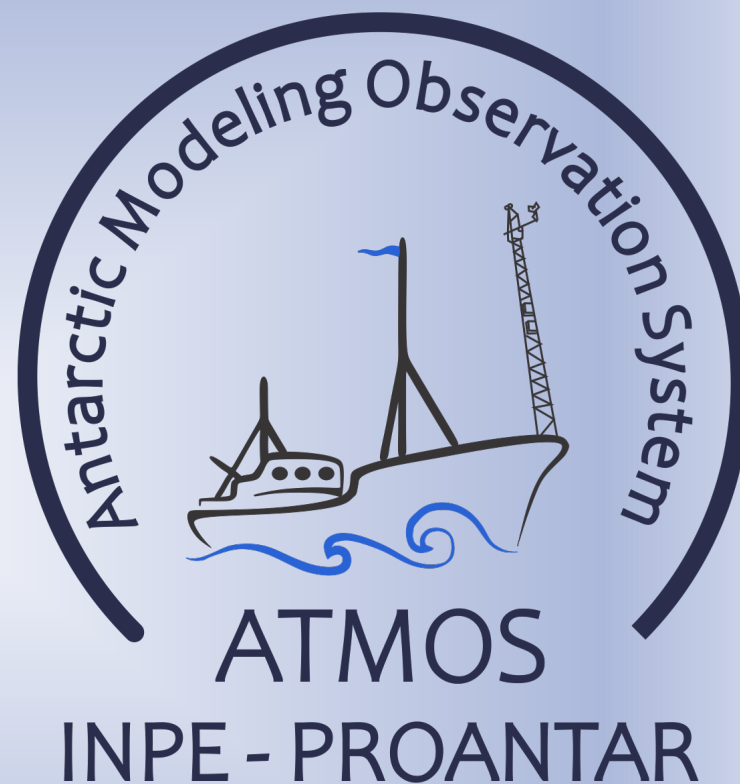
$$K = 0.36.U_{10m}^2 - 0.8 U_{10m} + 3.6$$

The **Warm Core Eddy** impact is perceived above the top of the Marine Atmospheric Boundary Layer (MABL) ...



a)

Thank you very much for the opportunity and your attention!!!



EUMETSAT-SOLAS
December - 2021